



Processing statistical parameters of concentration along a river network



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In-stream water quality modelling – state of the art

Driving forces of WQ modeling:

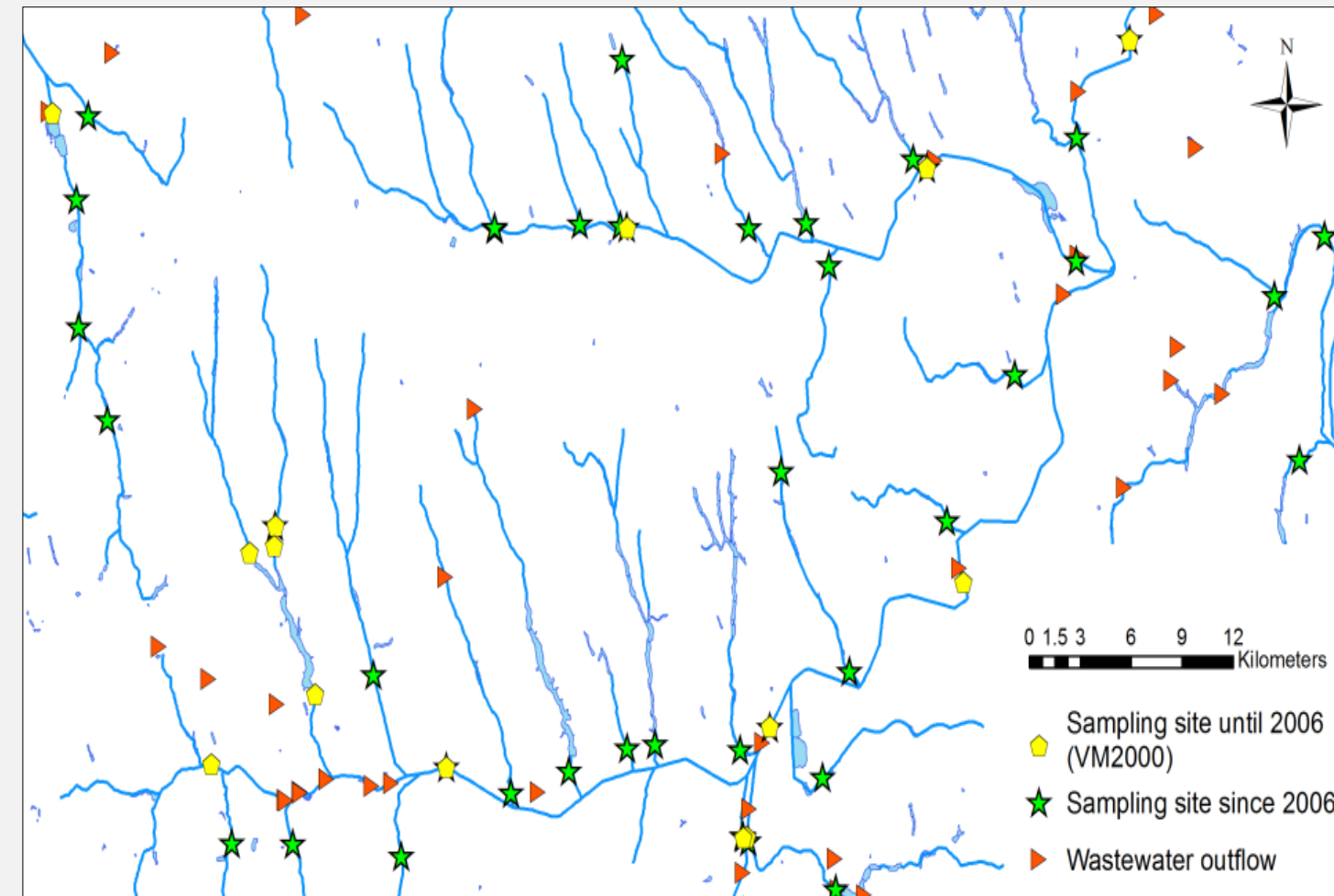
- contamination of waters (diffuse and point-sources)
- legislation (EU Directives, other international legislation,...)
- water uses (agricultural, domestic, industrial, recreational,...)
- scientific curiosity

Types of WQ models:

- the first step is always a hydrologic/ hydraulic model
- stochastic vs. process-based
- transient vs. stationary

Calibration/validation is problematic

- against extra measurements
- against another WQ model
- no calibration



Regular WQ monitoring in Hungary

Huge amount of data covering long period:
 ~ 50 years, > 1000 sampling sites, > 200 parameters

Time-shift between adjacent sampling stations

Data can be used for modelling purposes:

- long-term trends
- regional patterns

No use of the database to model ranges of concentrations in a longer time period with high spatial resolution!

Figure: Part of WQ monitoring network of Hungary

Modelling concept

- „steady-state” modelling of water quality in large spatial scale for „average” situation
- time variation accounted for with momentums (average, variance, covariances) of the time series
- make use of the regular monitoring data
- traditional parameters (BOD, DO, N, P)
- changes assigned to edges, mixing processes, diffuse and point-source loads and monit. points represented by nodes
- simplified process representation - second order second moment approximation of equations

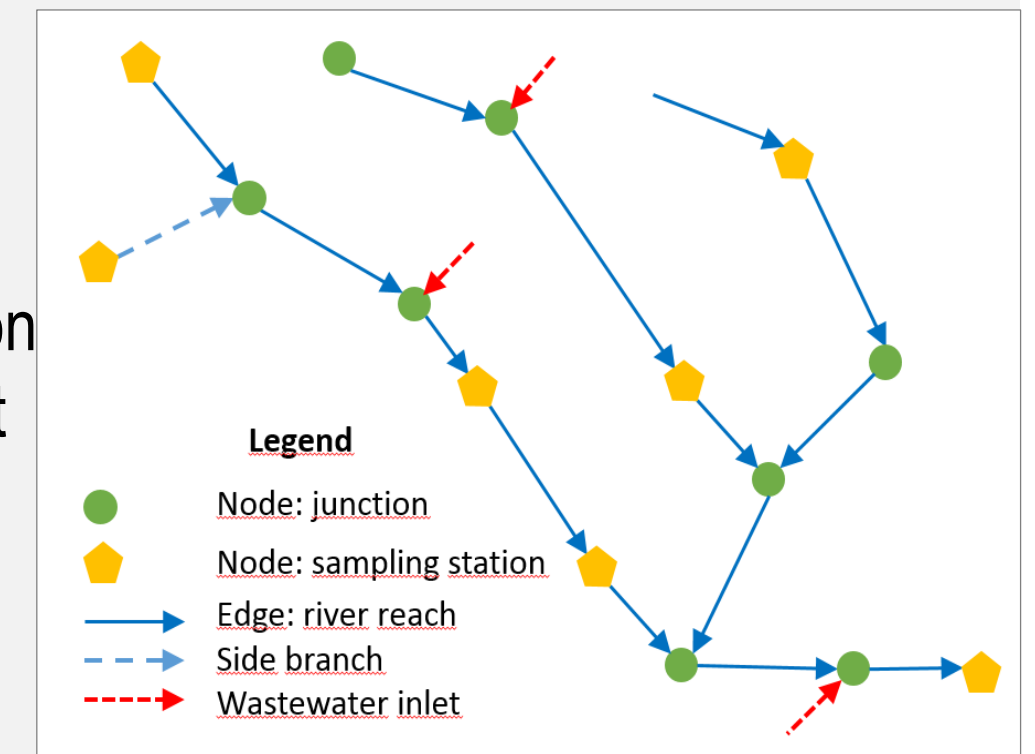


Figure: graph representation of the river network

Approximating the statistics

Edge process (settling):

$$C_1 = C_0 \exp\left(-\frac{v_s}{H} t\right)$$

where C_0 and C_1 – upstream and downstream concentration respectively, v_s – settling velocity, H – water depth and t – travel time along the river section.

$$C_1(C_0, Q) = C_0 \exp\left(k Q^{-\frac{5}{8}}\right)$$

where k comprises the settling velocity, the roughness and the geometrical properties and is assumed to be constant along an edge (a section of a reach).

Second-order approximation for mean downstream concentration:

$$E(C_1) \approx E(C_0) \exp\left(k [E(Q)]^{-\frac{5}{8}}\right) + \frac{1}{2} \left(\frac{\partial^2 C_1}{\partial C_0^2}\right)_{E(C_0), E(Q)} \cdot \sigma_{C_0}^2 + \left(\frac{\partial^2 C_1}{\partial C_0 \partial Q}\right)_{E(C_0), E(Q)} \cdot \text{cov}(C_0, Q)$$

where E – expected value (mean value); $\sigma_{C_0}^2$ – variance of water flow and $\text{cov}(C_0, Q)$ – covariance between upstream concentration and water flow.

Second-order approximation for standard deviation of downstream concentration:

$$\sigma^2(C_1) \approx \left[\bar{C}_1 \exp\left(\text{const}_2 Q^{-\frac{5}{8}}\right) - \bar{C}_1\right]^2 + \frac{1}{2} \left(\frac{\partial^2 f}{\partial C_0^2}\right)_{(\bar{C}_0, \bar{Q})} \cdot \sigma^2(C_0) + \frac{1}{2} \left(\frac{\partial^2 f}{\partial Q^2}\right)_{(\bar{C}_0, \bar{Q})} \cdot \sigma^2(Q) + \left(\frac{\partial^2 f}{\partial C_0 \partial Q}\right)_{(\bar{C}_0, \bar{Q})} \cdot \text{cov}(C_0, Q)$$

where $f = \left[C_0 \exp\left(\text{const}_2 Q^{-\frac{5}{8}}\right) - \bar{C}_1\right]^2$.

Second-order approximation for downstream correlation:

$$\text{cov}(Q, C_1) \approx \frac{1}{2} \left(\frac{\partial^2 g}{\partial C_0^2}\right)_{\bar{C}_0, \bar{Q}} \cdot \sigma^2(C_0) + \frac{1}{2} \left(\frac{\partial^2 g}{\partial Q^2}\right)_{\bar{C}_0, \bar{Q}} \cdot \sigma^2(Q) + \left(\frac{\partial^2 g}{\partial C_0 \partial Q}\right)_{\bar{C}_0, \bar{Q}} \cdot \text{cov}(C_0, Q)$$

where $g = \left(C_0 \exp\left(\text{const}_2 Q^{-\frac{5}{8}}\right) - \bar{C}_1\right) (Q - \bar{Q})$.

Node process (instant complete stirring):

$$C = \frac{Q_1 C_1 + Q_2 C_2}{Q_1 + Q_2} = h(Q_1, C_1, Q_2, C_2)$$

Mean of node outflow concentration:

$$E(C) = \frac{Q_1 \bar{C}_1 + Q_2 \bar{C}_2}{Q_1 + Q_2} + \frac{\bar{C}_1 - \bar{C}_2}{(Q_1 + Q_2)^{\frac{2}{3}}} (\bar{Q}_1 \sigma_{Q_2}^2 - \bar{Q}_2 \sigma_{Q_1}^2) + \frac{(\bar{C}_1 - \bar{C}_2)(\bar{Q}_1 - \bar{Q}_2)}{(Q_1 + Q_2)^3} \text{cov}(Q_1, Q_2) + \frac{\bar{Q}_2}{(Q_1 + Q_2)^2} \text{cov}(Q_1, C_1) + \frac{\bar{Q}_1}{(Q_1 + Q_2)^2} \text{cov}(Q_2, C_2) - \frac{\bar{Q}_2}{(Q_1 + Q_2)^2} \text{cov}(Q_1, C_2) - \frac{\bar{Q}_1}{(Q_1 + Q_2)^2} \text{cov}(Q_2, C_1)$$

Variance of node outflow concentration:

$$\sigma^2(C) = \left(\frac{Q_1 \bar{C}_1 + Q_2 \bar{C}_2}{Q_1 + Q_2} - \bar{C}\right)^2 + \left(\frac{\partial^2 h}{\partial Q_1^2}\right)_{\bar{C}_1, \bar{C}_2, \bar{Q}_1, \bar{Q}_2} \sigma_{Q_1}^2 + \left(\frac{\partial^2 h}{\partial Q_2^2}\right)_{\bar{C}_1, \bar{C}_2, \bar{Q}_1, \bar{Q}_2} \sigma_{Q_2}^2 + \left(\frac{\partial^2 h}{\partial C_1^2}\right)_{\bar{C}_1, \bar{C}_2, \bar{Q}_1, \bar{Q}_2} \sigma_{C_1}^2 + \left(\frac{\partial^2 h}{\partial C_2^2}\right)_{\bar{C}_1, \bar{C}_2, \bar{Q}_1, \bar{Q}_2} \sigma_{C_2}^2$$

Tests

Test 1: synthetically generated data

Input value	Range	Step	Units
Riverbank slope (φ)	15 ... 45	15	deg
Riverbed slope (S)	10^{-5} ... 10^{-4}	4.5	-
Manning's coefficient (k_{st})	10 ... 40	10	$m^{1/3} s^{-1}$
Settling velocity (v_s)	0.1 ... 10	*10	$m d^{-1}$
Mean discharge ($E(Q)$)	0.1 ... 2000	-	$m^3 s^{-1}$
Coefficient of variation for discharge ($\sigma(Q)/E(Q)$)	0.2 ... 1.5	-	-
Mean upstream concentration ($E(C_0)$)	10.0 ... 400	-	$g m^{-3}$
Coefficient of variation for concentration ($\sigma(C_0)/E(C_0)$)	0.5 ... 3.0	-	-
Discharge: parameter μ of logN random variables	-1 ... 7	1	$\ln(m^3 s^{-1})$
Discharge: parameter σ of logN random variables	0.2 ... 1	0.2	$\ln(m^3 s^{-1})$
Concentration: parameter μ of logN random variables	2 ... 5	1	$\ln(g m^{-3})$
Concentration: parameter σ of logN random variables	0.5 ... 1.5	0.5	$\ln(g m^{-3})$
Correlation coefficient $r(Q, C_0)$	-0.2 ... 0.8	0.2	-

Test 2: Measured upstream versus calculated downstream data

- 35-years long series of water flow and total suspended solids (TSS), measured daily on the Zala river at Zalaapáti between 1978 – 2012.
- 350 randomly selected, one-year long sections of measured discharge-concentration data pairs were selected.
- Settling velocity and river geometry were determined randomly in the same range as in Test 1

Input value	Range	Units
Reach length (L)	100	km
Riverbank slope (φ)	26.5	deg
Riverbed slope (S)	10^{-6}	-
Manning's coefficient (k_{st})	40	$m^{1/3} s^{-1}$
Settling velocity (v_s)	1	$m d^{-1}$

Test 3: Measured upstream and downstream data

- upstream and downstream of Kis-Balaton Water Protection System
- 1998 – 2006

Input value	Value / Range	Units
Riverbank slope (φ)	11.3	deg
Riverbed slope (S)	10^{-6}	-
Manning's coefficient (k_{st})	40	$m^{1/3} s^{-1}$
Settling velocity (v_s)	0.4 – 3.7	$m d^{-1}$
Mean discharge ($E(Q)$)	2.0 – 5.8	$m^3 s^{-1}$
Coefficient of variation (CV) of discharge ($\sigma(Q)/E(Q)$)	0.3 – 1.1	-
Mean upstream concentration ($E(C_0)$)	11 – 65	$g m^{-3}$
CV of concentration $\sigma(C_0)/E(C_0)$	0.9 – 3.0	-
Correlation coefficient $r(Q, C_0)$	-0.25 – 0.52	-
Observed mean downstream (ds) concentration ($E(C_1)$)	2.9 – 7.8	$g m^{-3}$
Observed ds CV of concentration ($\sigma(C_1)/E(C_1)$)	0.6 – 1.3	-

Results

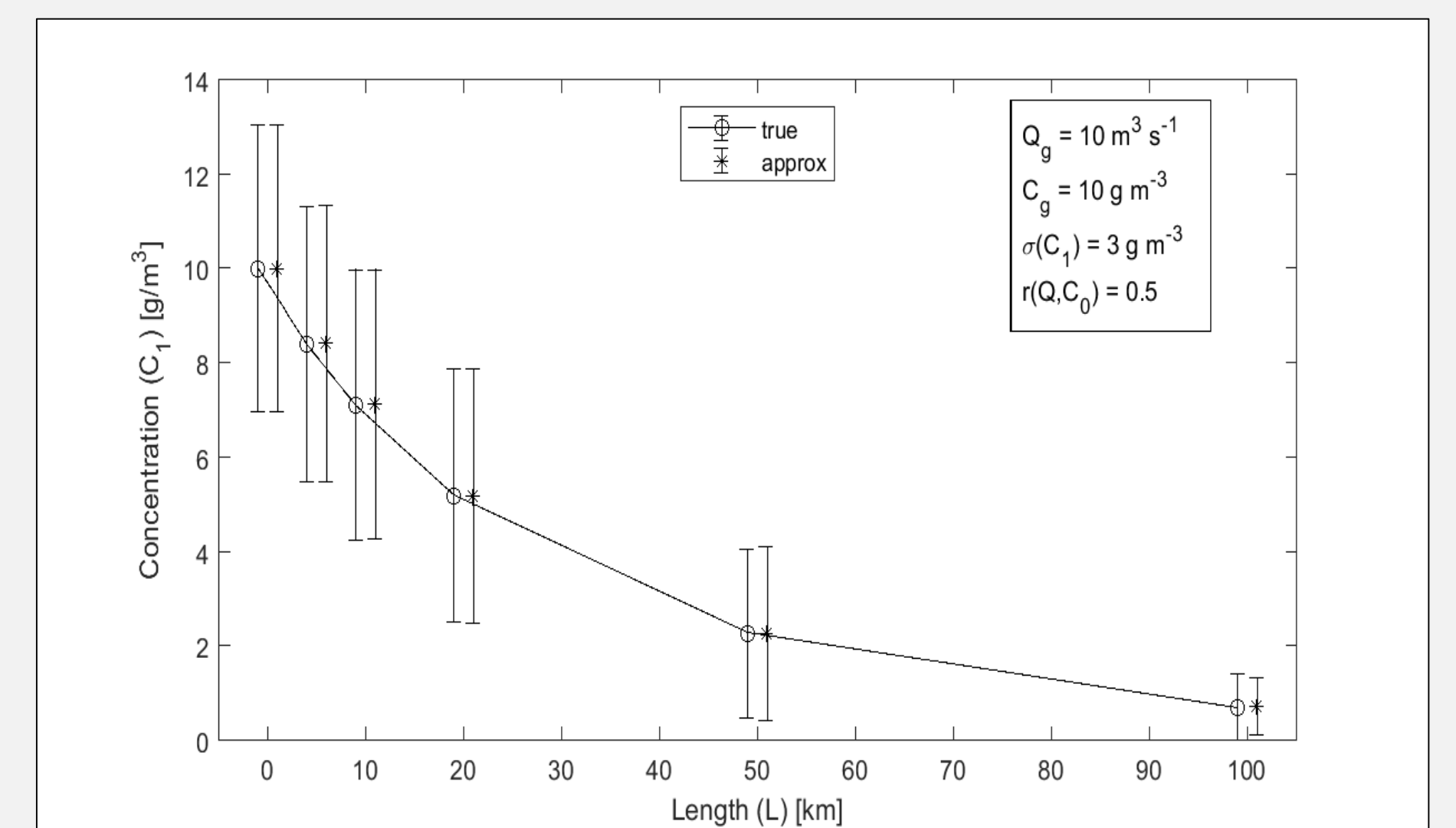


Figure: Longitudinal section of a unique test

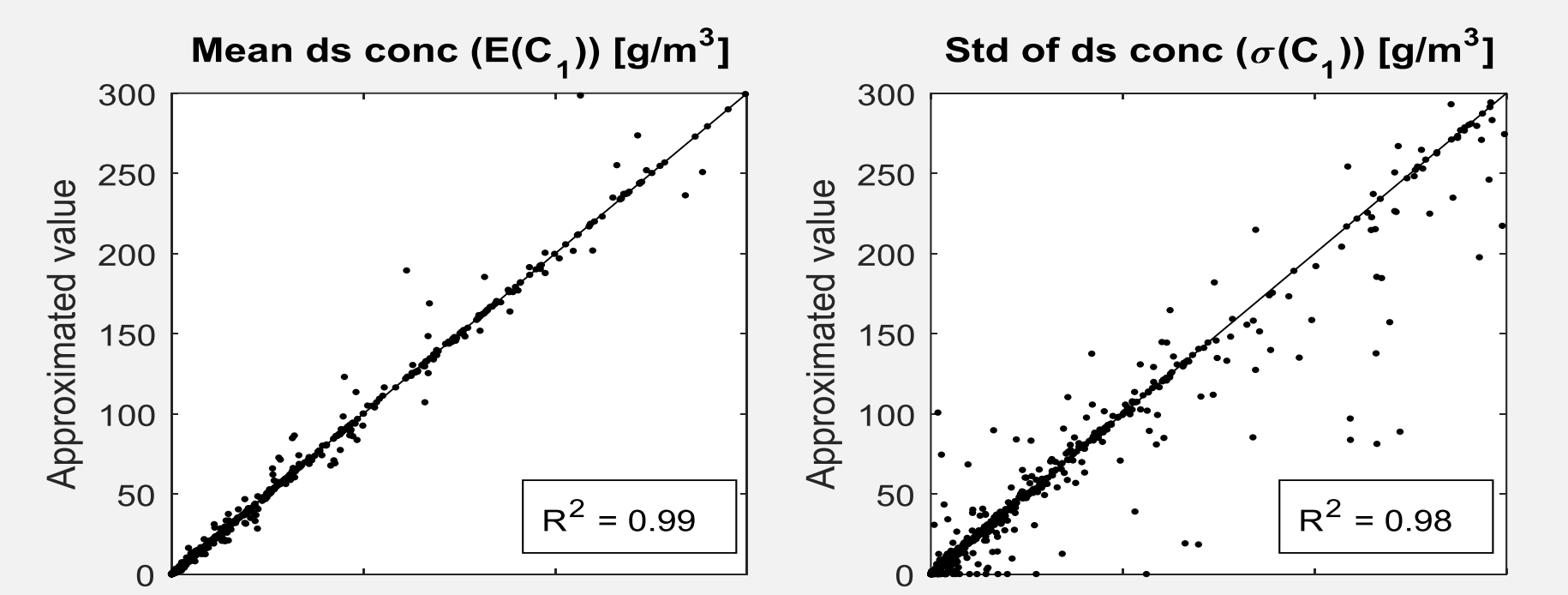


Figure: Approximated vs. true values

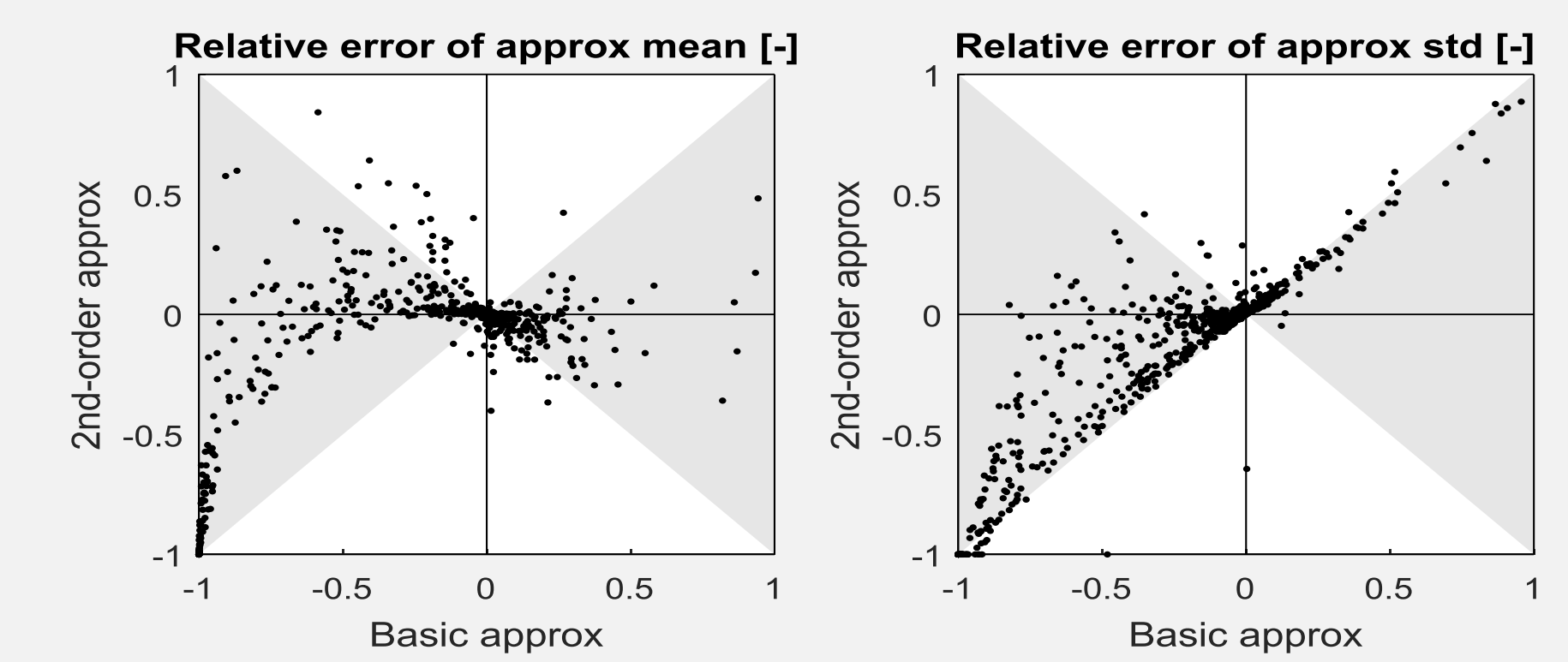


Figure: Errors of second-order approximations vs. errors of basic apx.

Results of Test 3:

year	Std of downstream concentration				Correlation between conc & disch					
	Obs [g/m ³]	Bapx [g/m ³]	Rerr of bapx	Apx2 [g/m ³]	Rerr of apx2	Obs ds [-]	Us [-]	Err of us	Apx2 ds [-]	Rerr of apx2
1998	8	12	51%	12	56%	0.11	0.23	0.11	0.72	0.61
1999	6	19	243%	11	98%	-0.04	0.28	0.33	0.56	0.60
2000	4	4	-13%	5	7%	-0.09	0.12	0.21	0.51	0.60
2001	8	6	-14%	7	-9%	0.08	0.52	0.43	0.64	0.56
2002	7	18	155%	19	182%	-0.47	-0.26	0.21	-0.19	0.27
2003	5	9	70%	11	110%	-0.19	-0.20	-0.01	0.06	0.25
2004	5	12	118%	9	66%	0.27	0.31	0.04	0.67	0.40
2005	2	8	259%	5	121%	0.03	0.20	0.17	0.62	0.59
2006	2	8	401%	3	97%	0.13	0.20	0.07	0.82	0.69

Conclusions

- The method can be considered as satisfactory in estimating downstream mean and variance.
- In almost all cases, relative error of the second-order approximation of mean downstream concentration remained very low.
- It performed well on both generated lognormal and measured data, compared to „basic” approximations.
- The approximated amounts were close to „true” values.
- Significant improvement in approximation of the mean downstream concentration

- In most cases, relative error of the second-order approximation of standard deviation of downstream concentration remained below 10%.
- Relative errors associated with smaller standard deviations sometimes remained high
- Approximation was unsatisfactory compared to downstream measurements, which indicates the need for more detailed description of the in-stream processes.

Next steps

- More detailed description of the in-stream process.
- Test on other WQ components
- Investigation of second-moment approximations of formulas more capable to describe complex in-stream processes
- Further investigation needed for
 - node processes
 - resuspension
 - diffuse pollution and/or continuous discharge-increment